

PRELIMINARY SENSITIVITY STUDY OF AEROSOL LAYER HEIGHT FROM SYNTHETIC MULTIANGLE POLARIMETRIC REMOTE SENSING MEASUREMENTS

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ABSTRACT:

Many previous studies have shown that multiangle, multispectral polarimetric remote sensing can provide valuable information on aerosol microphysical and optical properties, in which the aerosol layer height (ALH) is an important parameter but with less studies, especially in the near-ultraviolet (near-UV) and visible (VIS) wavelength bands. Based on the optimal estimation (OE) theory and information content analysis method, we focus on the sensitivity study of ALH with the synthetic data in the near-UV and VIS wavelength in the range of 410-865 nm, and further to assess the capability of multiangle intensity and polarization measurements for the retrieval of ALH. Unified Linearized Vector Radiative Transfer Model (UNL-VRTM) has been used as the forward model to simulate the intensity and polarized radiance at the top of atmosphere (TOA), as well as the Jacobians of TOA results with respective to corresponding parameters. The degree of freedom for signal (DFS) and a posteriori error are introduced to quantify the information content of ALH from the intensity and polarization measurements, respectively. By assuming the surface type, aerosol model, aerosol loads, prior errors and observation geometries, the sensitivity of ALH has been preliminarily investigated. The sensitivity study results show that the near-UV and polarization measurements are the important source of information content for the aerosol height retrieval in satellite remote sensing.

1. INTRODUCTION

Aerosol vertical distribution plays an important role in the study of radiative forcing, air quality, aerosol scattering in trace gas and ocean colour retrievals (Haywood and Boucher, 2000; Wang and Christopher, 2013; Butz et al., 2011; Gordon, 1997). For the active remote sensing, the aerosol vertical distribution can be effectively measured, such as the ground-based, airborne and space-borne measurements (Winker et al., 2010). While for the passive satellite remote sensing, the oxygen (O₂) A and B bands are usually used to study the aerosol vertical distribution based on synthetic data and real measurements (Sanghavi et al., 2012; Ding et al., 2014; Saders et al., 2015; Xu et al., 2017b, 2019). Many previous studies have shown that multiangle, multispectral polarimetric remote sensing can provide valuable information on aerosol microphysical and optical properties (Waquet et al., 2009; Hasekamp et al., 2011; Dubovik et al., 2011; Wu et al., 2015; Hou et al., 2018), in which the aerosol layer height (ALH) is one of important parameters but with less studies, especially in the near-ultraviolet (near-UV) and visible (VIS) wavelength bands (Wu et al., 2016).

The vertical distribution of aerosol is complex, which is not homogeneous due to the various structure depend on the height. The uncertainties from the treatment of the columnar atmosphere as homogeneous and independent as the height inevitably lead to the variations in aerosol optical features in model calculations (Bi et al., 2016). ALH is common used in the model calculation, while in the observational measurement, it is always instead of

planetary boundary layer height (PBLH). The planetary boundary layer is the lowest layer of the atmosphere, which contains the vast majority of aerosols of the low troposphere. ALH is an important factor of atmospheric diffusion ability, which determines the vertical disperse range of pollutants emitted from the ground (Wang et al., 2019).

As the flagship of the environment and atmosphere observation satellite in the Chinese High-resolution Earth Observation System (CHEOS) program (Gu and Tong, 2015), Gaofen-5 (GF-5) has been launched in May 2018, and the Directional Polarization Camera (DPC) is one of payloads (Li et al., 2018; Zheng et al., 2019). By inheritance the technology of Polarization and Directionality of the Earth's Reflectances (POLDER) (Deuzé et al., 2001), The DPC employed a charge coupled device (CCD) detection unit and realized spatial resolution of 3.3 km under a swath width of 1850 km, which has integrated 3 polarized channels (490, 670, 865nm) together with 5 non-polarized bands (443, 565, 763, 765, 910 nm). Besides, the Multi-Viewing-Channel-Polarization Imager (3MI) developed by EUMETSAT for the time-frame 2020-2040, will measure the radiance at the top of atmosphere (TOA) in 12 spectral channels, in which 9 channels will be polarized, including 410, 443, 490, 555, 670, 865, 1370, 1650, 2130 nm (Marbach et al., 2013).

In this paper, we focus on the preliminary sensitivity study of ALH with the synthetic data in the near-UV and VIS wavelength, and further to assess the capability of multiangle intensity and

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polarization measurements for the retrieval of ALH. For sensitivity study, the degree of freedom for signal (DFS) and a posteriori error of ALH have been employed. By assuming the surface type, aerosol model, aerosol loads, prior errors and multi-viewing observation geometries, the sensitivity of ALH could be systematically analysed and discussed. In section 2, we briefly describe the real aerosol vertical distribution based on the ground lidar measurements. Subsequently, we present the sensitivity study methodology in section 3, and then conduct the analysis of information content based on these synthetic measurements in section 4. Finally, we give the conclusions in section 5.

2. AEROSOL VERTICAL DISTRIBUTION

The common measurements of ALH are taken by radiosonde, Sodar, microwave radiometer and Lidar, of which lidar is the effective way to retrieve the continue variation of the ALH (Collaud et al., 2014; Li et al., 2017). The Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) has the capability to estimate the large-scale region ALH from space since 2006 (Winker et al., 2010). Meanwhile, similar measurements of the airborne Cloud Physics Lidar (CPL) has been utilized in many aircraft campaigns to profile the optical properties aerosol and cloud (McGill et al., 2002; Wu et al., 2016). Additionally, the ground-based lidar can provides more accurate measurements with high signal-to-noise ratio and time continuity (Wiegner et al., 2014; Tucker et al., 2009; Yang et al., 2013).

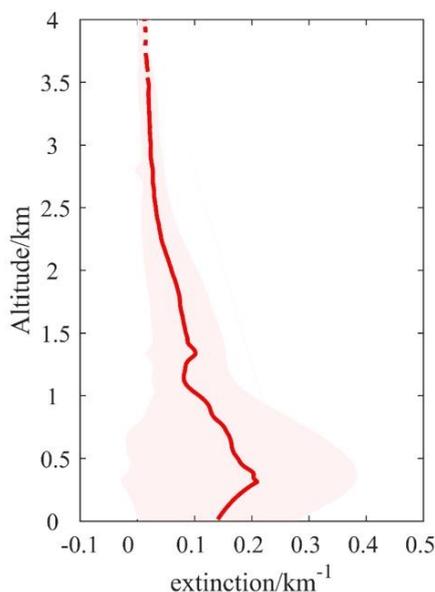


Figure 1. The mean extinction coefficients of aerosol in Oct. 6, 2017. The lines connected by red dots in the figure represent the average retrieved extinction coefficients, while the light red shadow indicates the standard deviation. The horizontal axis represents extinction coefficients and the vertical axis represents the attitude.

From the profile of aerosol extinction coefficients, we can clearly obtain the vertical distribution of aerosol. Based on the Fernald method, the aerosol extinction coefficient can be calculated (Fernald, 1984). Figure 1 shows the mean extinction coefficients of aerosol from a ground lidar called CE370 by CIMEL in France, which is located in the Institute of Remote Sensing and Digital Earth, Chinese Academy of Sciences in Beijing (Liu et al., 2017; Wang et al., 2019). We can see that the aerosol mainly exists in the height below 1 km, where the extinction coefficients grows

with a large gradient. At the height of 300 m, it seemly the concentration gets the peak value. However, it should be noted that due to the insufficient overlap, the uncertainty of measurement below 200m is large.

For lidar measurement, wavelet covariance transform is the most comment used method to identify ALH (Gamage and Hagelberg, 1993; Yang et al., 2013; Wang et al., 2019). The upper boundary of the AHL can be determined by the maximum variation gradient of the backscattering signal attenuated by aerosol (Seibert et al., 2000). From Figure 2, it is demonstrated that ALH starts to increase gradually from 08:00 a.m. and reaches its maximum at nearly 14: 00 p.m. From the midnight, ALH goes down rapidly and holds steady until the sunrise. Actually, the ALH shows a diurnal cycle from rising to decreasing along with the variation of the sunlight radiation.

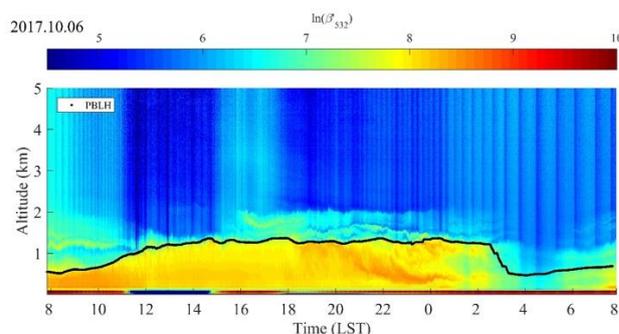


Figure 2. The aerosol layer height retrieved from wavelet covariance transform method in Beijing in Oct. 6, 2017. The lines connected by black dots in the figure represent the retrieved PBLH. The horizontal axis represents the local standard time (LST) and the vertical axis represents the height. Colorbar denotes the logarithm of the attenuated backscattering coefficient.

3. METHODOLOGY

3.1 DFS and Posterior Error

Following the optimal estimation (OE) theory (Rodgers, 2000; Dubovik et al., 2011), we have the forward model as

$$\mathbf{y} = \mathbf{F}(\mathbf{x}, \mathbf{b}) + \boldsymbol{\epsilon}, \quad (1)$$

where \mathbf{x} is a state vector to be retrieved, the vector \mathbf{b} contains these parameters that are not contained in \mathbf{x} but quantitatively influence the TOA measurements, \mathbf{y} is an observation vector, \mathbf{F} means a forward model, and $\boldsymbol{\epsilon}$ means an experimental error term integrated by the observation noise and forward model uncertainty.

For the sensitivity study, the averaging kernel matrix can be represented as

$$\mathbf{A} = (\mathbf{K}^T \mathbf{S}_\epsilon^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1} \mathbf{K}^T \mathbf{S}_\epsilon^{-1} \mathbf{K}, \quad (2)$$

where the superscript “ -1 ” is the inverse operation of matrix, \mathbf{S}_a means the error covariance matrix of the a priori estimate \mathbf{x}_a , the Jacobians matrix \mathbf{K} corresponds to the partial derivatives of $\mathbf{F}(\mathbf{x})$ with respect to \mathbf{x} . Besides, \mathbf{S}_ϵ is the covariance matrix of the error from both the measurements and the forward model with

$$\mathbf{S}_\epsilon = \mathbf{S}_y + \mathbf{K}_b \mathbf{S}_b \mathbf{K}_b^T, \quad (3)$$

in which \mathbf{S}_y is the measurement error covariance matrix, \mathbf{S}_b is the error covariance matrix for a vector \mathbf{b} of forward model, and \mathbf{K}_b is the Jacobians matrix of measurements \mathbf{y} with respect to the vector \mathbf{b} (Frankenberg et al., 2012; Xu and Wang, 2015; Hou et al., 2018a, b).

The DFS of each individual retrieved parameter is equal to $\mathbf{A}_{i,i}$, and in the range of 0 and 1. $\mathbf{A}_{i,i} = 1$ means that the observation is able to fully characterize the truth of x_i , and $\mathbf{A}_{i,i} = 0$ means that the observation do not contain any information on x_i at all. In other words, the closer the value of $\mathbf{A}_{i,i}$ to 1, the better the retrieval of parameter x_i . Meanwhile, a posteriori error covariance matrix $\hat{\mathbf{S}}$ can be written as

$$\hat{\mathbf{S}} = (\mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1}, \quad (4)$$

which means the statistical uncertainties in retrieved $\hat{\mathbf{x}}$ caused by measurement noise and the propagation of errors, and the posterior errors (absolute errors) are equal to the square root of diagonal elements (Hou et al., 2018a; Li et al., 2018).

3.2 Forward Simulations and Parameter Settings

To simulate the TOA measurements and corresponding Jacobian results in the near-UV and VIS wavelength from 410 nm to 865 nm of 3MI, Unified Linearized Vector Radiative Transfer Model (UNL-VRTM) is used as the forward model (Wang et al., 2014). UNL-VRTM integrates the VLIDORT code (Spurr et al., 2006), a linearized Mie code, a linearized T-Matrix code, a Rayleigh scattering module, line-by-line gas absorption calculation with HITRAN database, and some easy-used surface modules including the bidirectional reflectance distribution function (BRDF) and bidirectional polarized reflectance distribution function (BPDF). Based on UNL-VRTM, many studies have been carried to analyse the information content and develop the inversion algorithm for retrievals from the various measurements (Chen et al., 2017; Ding et al., 2016; Hou et al., 2016, 2017, 2018a, b; Xu and Wang, 2015; Xu et al., 2015, 2017a, b, 2018, 2019; Li et al., 2018; Zheng et al., 2019a, 2019b).

The kernel-driven BRDF model used in UNL-VRTM for surface reflectance can be written as

$$\rho_s(\mu_0, \mu_v, \phi, \lambda) = f_{iso}(\lambda) + k_1(\lambda) f_{geom}(\mu_0, \mu_v, \phi) + k_2(\lambda) f_{vol}(\mu_0, \mu_v, \phi), \quad (5)$$

where μ_0 and μ_v respectively represent the cosine of solar zenith angles and viewing zenith angle; f_{iso} , f_{geom} and f_{vol} respectively corresponds to the isotropic, geometric-optical (Li-sparse kernel) and volumetric (Ross-thick kernel) surface scattering, $f_{iso}(\lambda)$; $k_1(\lambda)$ and $k_2(\lambda)$ are the coefficients of the BRDF kernels at the wavelength λ (Litvinov et al., 2011). For the surface polarized reflectance, the used BPDF is

$$R_p^s(\mu_0, \mu_v, \varphi) = \frac{C \exp(\tan \gamma) \exp(\text{NDVI})}{4(\mu_0 + \mu_v)} \mathbf{F}_{1,2}(\gamma, n_i), \quad (6)$$

where $\mathbf{F}_{1,2}(\gamma, n_i)$ represent the polarized component of the Fresnel reflection matrix, n_i is the reflective index of the vegetative matter, γ means the half the phase angle, NDVI is the normalized difference vegetation index, and C is the only free linear parameter (Maignan et al., 2009). In this paper, he parameters of the BRDF and BPDF of vegetated surface are chosen from the work of Litvinov et al. (2011), which will not be discussed here anymore (Hou et al., 2019).

No.	Used bands combination
1	410 nm
2	410, 443 nm
3	410, 443, 490 nm
4	410, 443, 490, 555 nm
5	410, 443, 490, 555, 670 nm
6	410, 443, 490, 555, 670, 865 nm

Table 1. Number of used bands combination for sensitivity study of ALH with the multi-viewing intensity and polarization measurements

In order to obtain the synthetic data of multiangle for sensitivity study, 2 typical multi-viewing observation geometries are considered with the combinations of different solar zenith angles (θ_0), viewing zenith angles (θ_v) and relative azimuth angles (ϕ) to represent the observations in different location. Figure 3 illustrates the polar-plot observational geometries, in which the radius corresponds to the θ_v change from 0° to 60° with the step of 20° , and the polar angle means ϕ change from 0° to 360° , as well as the position of Sun also defined with $\phi = 0^\circ$ (Zheng et al., 2019a). Here, the geometry $\phi = 0^\circ$ means the observer and Sun are in the same direction and also in the same side of main plane, while $\phi = 180^\circ$ represents the opposite direction and side of main plane, so as the definition in other polar-plot figures.

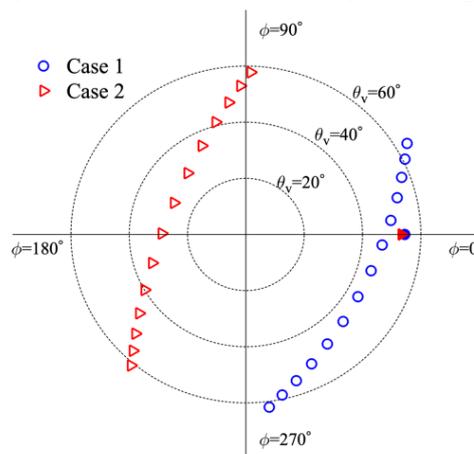


Figure 3. Multi-viewing observation geometries for forward simulations for 2 different cases. In which the solid circle and triangle with $\phi = 0^\circ$ represent the corresponding position of the Sun for case 1 and case 2, respectively.

For the sensitivity study of ALH, Table 1 lists the number of used bands combination with the multi-viewing intensity and polarization measurements, in which No.1 means only the 410 nm wavelength is used, and then add the next VIS wavelength in sequence until all of the 6 wavelength bands are used from 410 to 865 nm. Figure 4 illustrates the angular distribution of BRDF results in 410 nm and the angular distribution of wavelength-independent BPDF results for the vegetated surface. The angular distribution of BRDF results for other bands are not shown here.

For the aerosol vertical profile parameter settings in UNL-VRTM, the exponentially decreasing profile with scale height H given is used (Wang et al., 2014; Xu and Wang, 2017), which can be written in the form as

$$\int_{TOA}^z \tau(z) dz = \tau_0 \exp\left(-\frac{z}{H}\right), \quad (7)$$

where τ_0 is the columnar aerosol optical depth (AOD), and we set the scale height H is equal to 2 km in this study. Besides, we

set the other aerosol parameters of particle size distribution and complex refractive index with fine-dominated aerosols and AOD = 0.8 at 550nm by following the work of Hou et al. (2017, 2018).

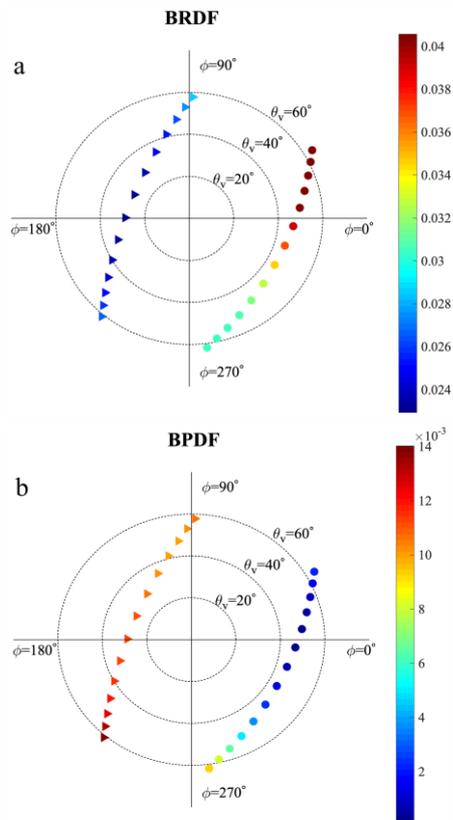


Figure 4. The angular distribution of BRDF results in 410 nm (a) and the angular distribution of wavelength-independent BPDF results (b) with each colour bar for vegetated surface.

For the information content analysis of ALH in the near-UV and VIS wavelength bands, the state vector only contains one parameter, that is

$$\mathbf{x} = [H]^T, \quad (8)$$

with the vector

$$\mathbf{b} = [V_0^f, V_0^c, \rho_s(\lambda)]^T, \quad (9)$$

where V_0^f and V_0^c represents the fine-mode and coarse-mode columnar volume concentration respectively, $\rho_s(\lambda)$ is the surface BRDF at the wavelength λ . The measurement vector \mathbf{y} can be defined as

$$\mathbf{y} = [I_\lambda^{v_1}, \dots, I_\lambda^{v_M}, Q_\lambda^{v_1}, \dots, Q_\lambda^{v_M}, U_\lambda^{v_1}, \dots, U_\lambda^{v_M}]^T, \quad (10)$$

in which, I , Q and U are the TOA measured Stokes elements, the superscript $v_1 - v_M$ are used to note each viewing in multiple observation. Correspondingly, the Jacobian vector and matrix

$$\mathbf{K} = \left[\frac{\partial I_\lambda^{v_1}}{\partial H}, \dots, \frac{\partial I_\lambda^{v_M}}{\partial H}, \frac{\partial Q_\lambda^{v_1}}{\partial H}, \dots, \frac{\partial Q_\lambda^{v_M}}{\partial H}, \frac{\partial U_\lambda^{v_1}}{\partial H}, \dots, \frac{\partial U_\lambda^{v_M}}{\partial H} \right]^T, \quad (11)$$

with

$$\mathbf{K}_b = \begin{bmatrix} \frac{\partial I_\lambda^{v_1}}{\partial V_0^f} & \frac{\partial I_\lambda^{v_1}}{\partial V_0^c} & \frac{\partial I_\lambda^{v_1}}{\partial \rho_s(\lambda)} \\ \vdots & \vdots & \vdots \\ \frac{\partial I_\lambda^{v_M}}{\partial V_0^f} & \frac{\partial I_\lambda^{v_M}}{\partial V_0^c} & \frac{\partial I_\lambda^{v_M}}{\partial \rho_s(\lambda)} \end{bmatrix}. \quad (12)$$

For the error covariance matrix

$$\mathbf{S}_a = [\sigma_H^2], \quad (13)$$

here we set $\sigma_H = 75\%H$ with 1.5 km, and

$$\mathbf{S}_b = \text{diag} \left([(\Delta V_0^f)^2, (\Delta V_0^c)^2, (\Delta \rho_s(\lambda))^2]^T \right), \quad (14)$$

with prior error of 20% for V_0^f and V_0^c , and 10% for $\rho_s(\lambda)$. For the measurement error covariance matrix, we have

$$\mathbf{S}_y = \text{diag} \left([(\Delta I_\lambda^{v_1})^2, \dots, (\Delta I_\lambda^{v_M})^2, \dots, (\Delta U_\lambda^{v_1})^2, \dots, (\Delta U_\lambda^{v_M})^2]^T \right). \quad (15)$$

By considering the integrated effects of instrumental noise with radiometric calibration and polarimetric accuracy uncertainties, we assume that the errors of measurements are 5% for both the intensity and the polarization for the sensitivity study (Li et al., 2018).

4. RESULTS

4.1 Simulated TOA Results

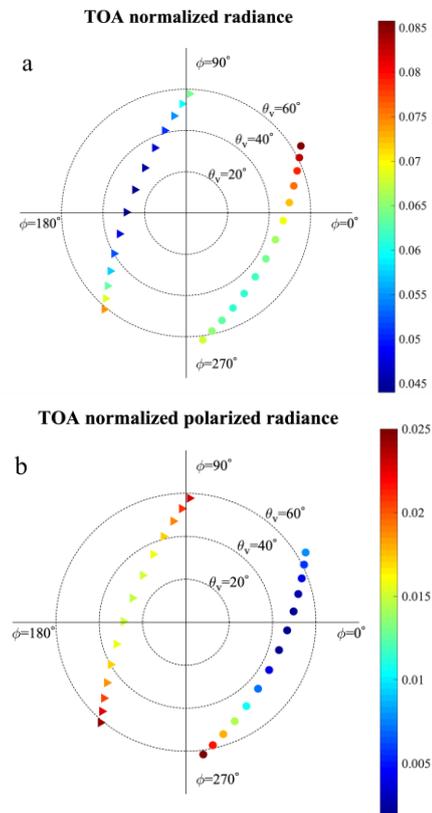


Figure 5. The angular distribution of TOA normalized radiance (a) and normalized polarized radiance (b) in 410 nm with vegetated surface and AOD = 0.8 at 550 nm.

To obtain the synthetic data for the sensitivity study of ALH, 2 multi-viewing observation cases are considered, and the TOA synthetic data are calculated using UNL-VRM with the vegetated surface and fine-dominated model, in which the aerosols are mixed by the fine- and coarse-mode particles with the fine-mode fraction of columnar volume concentration $FMF_V = 0.8$ (Hou et al., 2017, 2018). Figure 5 shows the angular distribution of TOA normalized radiance and polarized radiance in 410 nm. Here, the normalized radiance is equal to the first Stokes element I , while the normalized polarized radiance is calculated by $\sqrt{Q^2 + U^2}$. Figure 6 shows the angular distribution of Jacobians of TOA Stokes elements (I , Q and U) with respect to the aerosol parameters V_0^f and H in 410 nm, respectively. Due to limited space, the angular results in other wavelengths are not shown here.

4.2 Sensitivity Study Results

For the DFS and posterior errors of each used bands combination listed in Table 1, we consider the multi-viewing measurements of only intensity and combined with polarization together respectively, and then calculate the mean results of 2 multi-viewing observation cases as the final results.

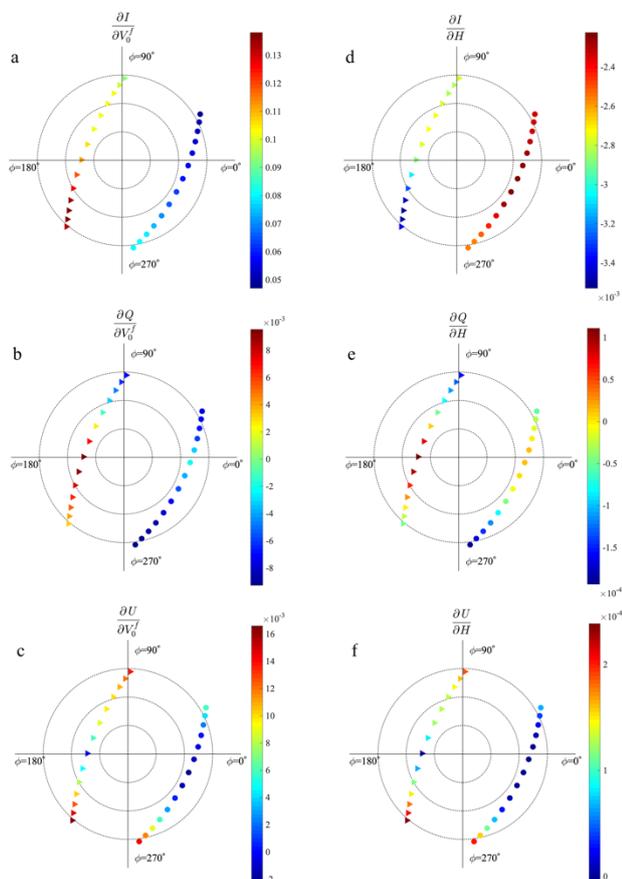


Figure 6. The angular distribution of Jacobians of TOA Stokes elements I , Q and U with respect to V_0^f (a-c) and H (d-f) in 410 nm.

Figure 7 shows the DPS of ALH along the number of used bands presented in the format of histogram. For the measurements of intensity, the DFS of ALH increases from 0.59 to 0.62 with the increasing number of used bands from 1 to 6 in the range of 410-865nm in sequence. While combining the measurements of polarization with intensity together, the information content of

ALH has a significant improvement, and the DFS of ALH increase from 0.94 to 0.98.

Figure 8 illustrates the posterior errors of ALH along the number of used bands with 75% priori errors of AHL (absolute priori errors 1.5 km). When only the intensity measurements are used, the absolute posterior error decreases from 0.64 km to 0.61 km in the range of 410-865 nm in in sequence. After the polarization measurements are added, the absolute posterior error can decrease from 0.38 km to 0.21 km. That is to say, the polarization can provide more useful information and constraints for the retrieval of ALH from multi-viewing measurements.

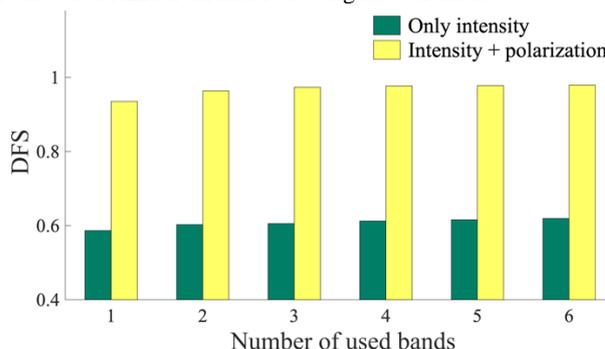


Figure 7. The DPS of ALH along the number of used bands for “only intensity” and “intensity + polarization” multi-viewing measurements.

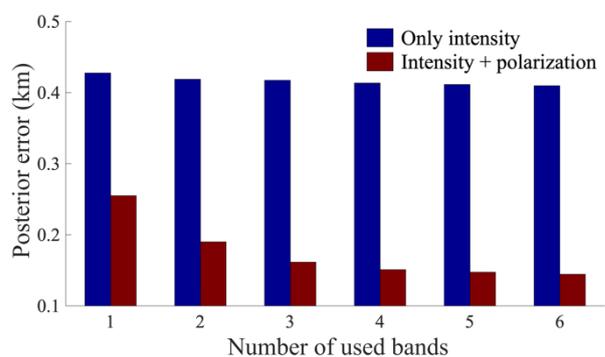


Figure 8. The posterior errors of ALH along the number of used bands for “only intensity” and “intensity + polarization” multi-viewing measurements with 75% priori errors of ALH.

5. CONCLUSIONS

In the paper, based on the OE theory and information content analysis method, the sensitivity study of aerosol layer height is carried out with the synthetic data in the range of 410-865 nm, and the capability of multiangle intensity and polarization measurements for the retrieval of ALH has been preliminarily assessed. UNL-VRM is used as the forward model to calculate the normalized radiance and polarized radiance at the top of atmosphere (TOA), as well as the Jacobians of TOA results with respect to corresponding parameters. The degree of freedom for signal (DFS) and a posteriori error are introduced to quantify the information content of ALH from the intensity and polarization measurements, respectively. By assuming the vegetated surface type, fine-dominated aerosol model, aerosol loads, prior errors and 2 typical multiple observation geometry cases, the preliminarily sensitivity study results show that the near-UV and polarization measurements are the important source of information content for the aerosol height retrieval in satellite remote sensing.

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